

Enrichment of redox-sensitive trace metals (U, V, Mo, As) associated with the late Hauterivian Faraoni oceanic anoxic event

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Abstract The Faraoni Level is a short-lived oxygen-deficient event that took place during the latest Hauterivian. In order to improve our understanding of the palaeoenvironmental conditions that occurred during this event, we have analysed the contents of several redox-sensitive trace elements (U, V, Mo, As, Co, Cd, Cu, Zn, Ni, Pb, Cr) from bulk limestone samples of late Hauterivian–early Barremian age from three reference sections. U, V, Mo and As show consistent and significant enrichments during the Faraoni event whereas the other redox-sensitive trace elements analysed here are not systematically enriched. In order to explain this discrepant behaviour, we propose that the Faraoni Level was deposited during a period of anoxic conditions near the sediment–water interface. The distinctive peaks in U, V, Mo and As contents are traceable throughout the three studied sections and represent a good correlation tool which helps to identify the Faraoni Level and its equivalents in the western

Tethyan realm and outside of the Tethys. For example, a peak in U contents in upper Hauterivian sediments of the northwestern Pacific realm (ODP leg 185, site 1149) may well be an expression of the Faraoni event in this particular basin.

Keywords Hauterivian · Barremian · Chemostratigraphy · Western Tethys · Trace metals

Introduction

Since the first description by Cecca et al. (1994) of an organic-rich guide level—the so-called Faraoni Level—in upper Hauterivian (late *Spathicrioceras angulicostatum* ammonite zone) pelagic carbonates of the Umbria-Marche Basin (Italy), it has been shown that correlatable organic-rich levels occur throughout the western Tethyan realm. Equivalents of the Faraoni Level have been identified on the Trento Plateau in Italy (Cecca et al. 1996; Coccioni et al. 1998), in the Vocontian Trough in France (Baudin et al. 1999), and in the Swiss Préalpes (Busnardo et al. 2003). It is for this reason that Baudin et al. (2002) associated this level with a short-lived oxygen-deficient event in the western Tethyan realm. And following Baudin (2005), the Faraoni Level and its equivalent can be qualified as an anoxic event (compare this reference for a detailed biostratigraphic definition of this event). With the Valanginian event (Erba et al. 2004), the Faraoni event is one of the first palaeoceanographic events which led to major change in the carbon cycle during the early Cretaceous.

The environmental circumstances leading to this short-lived oxygen-deficient event in the late Hauterivian are still not clear. In contrast to the anoxic events

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in the Valanginian, the early Aptian, and near the Cenomanian/Turonian boundary (e.g. Schlanger and Jenkyns 1976; Schlanger et al. 1987; Lini et al. 1992; Jones and Jenkyns 2001; Sanfourche and Baudin 2001), the Faraoni Level is not accompanied by a major carbon isotope excursion toward more positive values, but rather by a small long-term increase culminating around the Faraoni event (Föllmi et al. 1994; Erba et al. 1999; Van de Schootbrugge et al. 2000; Company et al. 2005; Godet et al. 2006). Furthermore, no major volcanic episode has, of yet, been identified for the late Hauterivian. This means that the classical model postulated for the origin of oceanic anoxic events in general, i.e. increased volcanism leading to an increase in atmospheric CO₂, climate warming, increased weathering and nutrient availability, higher productivity and consequently higher preservation rates of organic matter (e.g. Jenkyns 1999, 2003) cannot be applied in this case.

The Faraoni event corresponds in time to the onset of drowning episode D3 during the evolution of a carbonate platform system located along the northern Tethyan margin (Föllmi et al. 1994). But, in contrast to the Valanginian and the Aptian–Albian drowning episodes (D1, D4 and D5), the drowning episode D3—not being mirrored by a distinct positive excursion in carbon isotopes, is also in need of an alternative explanation (Van de Schootbrugge et al. 2000, 2003). It is therefore important to identify the palaeoenvironmental conditions leading to the Faraoni event, whilst also respecting the coeval drowning event on the northern Tethyan carbonate platform, which may have been linked to this anoxic event.

We investigated three (hemi-) pelagic key sections representing a north–south transect through the western Tethys and including micritic carbonates and marls of the Hauterivian–Barremian transition with regard to their redox-sensitive trace-element concentrations. The studied trace elements are uranium (U), vanadium (V), molybdenum (Mo), arsenic (As), cobalt (Co), cadmium (Cd), copper (Cu), zinc (Zn), nickel (Ni), lead (Pb) and chromium (Cr). Their behaviour around the Faraoni Level is quite distinct and helps us to improve our understanding of redox conditions and changes therein in this basin during the Faraoni event.

In this paper, the term “Faraoni Level” and “Faraoni Level equivalent” refers to the interval composed of organic-rich sediments and the intervening carbonate beds that are associated with the Faraoni oceanic anoxic event. The term “Faraoni Level” is used when the lithological characteristics are the same as in the type area (e.g. Cecca et al. 1994), whereas the term “Faraoni Level equivalent” is used when the

lithology is different from that of the type area (i.e. lacking a distinct enrichment in organic matter). The redox classification of sedimentary environments used here is adapted from Algeo and Maynard (2004) and references therein. Thus, oxic, suboxic and anoxic environments are characterized by >2.0, 2.0–0.2 and <0.2 ml O₂ l⁻¹H₂O, respectively. Anoxia and euxinia are distinguished by the absence or presence of free H₂S in the water column, respectively.

Studied sections and palaeogeographic setting

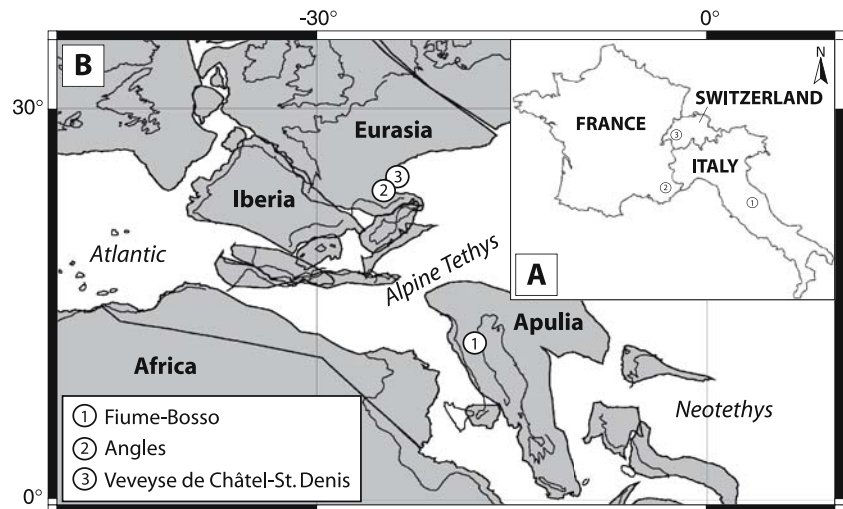
The studied sections were chosen according to the following criteria: (1) presence of the Faraoni Level or its equivalent; (2) good age control based on biostratigraphy (ammonites, nannofossils); (3) presence of pelagic carbonates; and (4) representative of different areas of the western Tethys (Fig. 1).

The first studied section is the Fiume-Bosso section, located between Urbino and Gubbio, near Cagli (central Italy; see also Cecca et al. 1994). This section represents the type section for the Faraoni Level (Cecca et al. 1994). The lithology consists of pelagic limestone and chert, which locally includes thin organic-rich marl layers and laminae, characteristic of the Maiolica Formation. The Faraoni Level is composed by a succession of three laminated black-shale layers and intervening calcareous beds.

The Veveysse de Châtel-St Denis (VCD) is the second studied section. It is situated in the canton of Fribourg, along the Veveysse river, in western Switzerland (see Busnardo et al. 2003, for a detailed geographic description). Its lithology consists of a succession of alternating pelagic marls and marly limestones, which are rich in macrofossils. The equivalent of the Faraoni Level (identified by ammonite dating) is represented by the “couches à poissons” (an alternation of more or less organic-rich shale and marly limestone, rich in fish fossils) (Busnardo et al. 2003).

The third studied section is along the “route d’Angles,” near Barrême in southeastern France, and is situated in the Vocontian Trough (see Busnardo 1965, for a detailed geographic description). Its lithology consists of a fossiliferous, hemi-pelagic succession of regularly alternating marlstone and limestone. Ammonite dating indicates that the Faraoni Level equivalent is situated around the bed AN 53. Black shale has not been distinguished within this interval, which is characterized by an alternation of limestone and laminated shale. In the sedimentary interval attributed to the *P. mortilleti* zone, no marls were sampled due to poor outcrop quality.

Fig. 1 **a** Actual location of the three studied sections. **b** Paleogeographic position of the three studied sections at the Hauterivian–Barremian boundary. Palaeomap modified from Hay et al. (1999). Modified from Bodin et al. (2006)



During the late Hauterivian, the Angles and the VCD sections were situated along the northern Tethyan margin. The VCD section is part of the Ultrahelvetic realm, which is considered as the deeper offshore prolongation of the Vocontian Trough to the northeast (Trümpy 1960). The Fiume-Bosso section, which is the deepest section, was situated in the southern part of the Tethys, in the Umbria-Marche Basin, remote from any continent.

Methods

ICP-MS analyses were performed on bulk rock samples from limestone and marlstone for all three sections. The limestone samples were sawed in order to eliminate altered parts and rock veins. Then, powders were obtained for both limestone and marls using a mechanic agate crusher. A portion of approximately 250 mg was transferred into a digestion vessel and 10 ml of concentrated nitric acid (suprapur, Merck, Darmstadt, Germany) were added. The sample was digested in a microwave oven (MSL-Ethos plus, Milestone) using the heating program recommended by the EPA 3051 method. After cooling, the resulting solution was filtered (0.45 μm) and diluted to 100 ml with ultrapure water. For the limestone, dissolution percentages determined after filtration were about 91% of initial sample weight in the Angles section, 83% in the VCD section, and 98% in the Fiume-Bosso section. Moreover, no correlation was observed between the concentration of the different analysed samples and the dissolution percentage obtained during digestion procedure (e.g. U in Fig. 2). This shows that the studied elements are present in the soluble

authigenic phase and are not due to partial leaching of the detrital insoluble fraction.

In the marls, the dissolution percentages were close to 50%. Mineralogical analyses indicate that clay and quartz grains constitute the insoluble part of the rock. No trace of carbonate is found, allowing us to postulate that the entire authigenic portion of the rock is dissolved with this method.

A second dilution (1/20) was then performed prior to analysis. To correct for matrix-induced ion signal variation and instrumental drift, rhodium was used as the internal standard. The element concentrations of the acid digests were determined by ICP-MS (Perkin-Elmer ELAN 6100, Perkin-Elmer Corp., Wellesley, MA, USA) using a semi-quantitative mode (Totalquant). Totalquant (TQ) is a simple, rapid and accurate panoramic method based on full mass-spectra scan methods and was successfully applied to the analysis of various biological and environmental samples (Jitaru et al. 2003) like, for instance, in sediments (Bayon et al. 1998). TQ enables spectral interpretation of a full mass spectrum by comparison with an internal response table which includes assignment of element intensities based on internal response factors and an interpretation of interferences based on implemented algorithms for numerical calculation of interference corrections. The response table is usually updated before analysis by determination of a blank and one standard solution. Two Certified Reference Materials (CRM) (NIST-1640 natural water, LKSD-1 lake sediment from Canadian Centre for Mineral and Energy Technology—CANMET—Canada) were measured to evaluate the accuracy which is achieved in panoramic analysis with the Elan 6100. Here we present the results obtained for the four principal trace elements

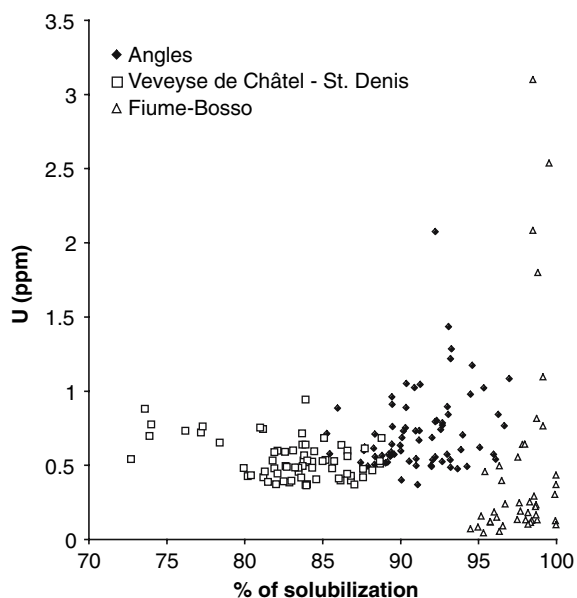


Fig. 2 U (ppm) versus percentage of dissolution in the three studied sections. No correlation is observed. See text for discussion

discussed in this study (V, Mo, As and U). The mean recovery rates were determined for V, Mo and As (no certified concentration values for U are available for the two CRM) and were, for the NIST-1640 CRM ($N = 11$) 101% for V, 99% for Mo and 101% for As, and relative standard deviation (RSD) was better than $\pm 9\%$ for V, $\pm 6\%$ for Mo and $\pm 7\%$ for As; for the LKSD-1 CRM ($N = 12$ digestions) mean recovery rates were 90% for V, 89% for Mo and 110% for As (for this CRM, recovery rates are based on certified concentration values obtained by partial extraction using concentrated HNO_3 and HCl whereas we used concentrated HNO_3 digestion); analytical precision (RSD) was about $\pm 8\%$ for V, $\pm 9\%$ for Mo and $\pm 12\%$ for As.

Furthermore, multiple digestions ($N = 10$) of a sample corresponding to the Faraoni Level of the Angles section (bed AN 53-2) were analysed. The RSD of these analyses were about $\pm 15\%$ for V, $\pm 12\%$ for Mo, $\pm 12\%$ for As and $\pm 3\%$ for U. A detection limit lower than $1 \mu\text{g}/\text{kg}$ was observed for the four trace elements studied.

The content and the type of organic matter were determined using Rock-Eval pyrolysis (Rock-Eval 6; Behar et al. 2001). The parameters TOC (total organic carbon in weight %) was obtained using the standard temperature cycle. Two standards (an in-house standard and the "IFP 160000" standard from the Institut Français du Pétrole, Paris, France) were passed at the beginning and at the end of a batch of approximately 15 samples.

Results

Redox-sensitive trace elements

The data for U, V, Mo, As, Co, Cd, Cu, Zn, Ni, Pb and Cr are shown in Tables 1 and 2 for limestone and marl for the three sections.

The Fiume-Bosso section

In limestone samples, the data obtained for U, V, Mo and As contents show comparable variations, with a consistently low background level, contrasted by a maximum in concentrations in and close to the Faraoni Level (Fig. 3; Table 1). The average values for the background level correspond to approximately 0.2, 0.9, 0.03 and 0.3 ppm and the maximum values to approximately 3.1, 11.6, 0.5 and 3.0 ppm, for the U, V, Mo and As contents, respectively.

The data obtained in the marlstones show equally a clear enrichment of U, V, Mo and As around the Faraoni Level (Table 2). The maximum values reach 15.4 ppm for U, 345.9 ppm for V, 61.8 ppm for Mo and 177.4 ppm for As, whereas background level are equal to approximately 2.9, 54.1, 2.3 and 13.2 ppm, respectively.

Despite a small negative shift in the upper part of the section, the global trend of Co in limestone samples seems to be more or less constant (Table 1), fluctuating around 1.8 ppm. The Faraoni Level is marked by a short negative shift toward 0.9 ppm. In the shale, the Co curve shows a small positive shift within the Faraoni Level. A maximum of 106.6 ppm is present just above the Faraoni Level (Table 2) whereas the background level values correspond to approximately 50 ppm.

For the other elements, only Cd and Zn show a small positive shift in limestone levels within the Faraoni Level. In the upper part of the section, a positive shift (up to 66.1 ppm) of Zn is remarked. In the Faraoni shale samples, a positive shift of Cd, Cu, Zn, Ni is present, with very high maximum values (10.5, 273.7, 1068.2 and 533.5 ppm, respectively). On the other hand, it appears that Pb and Cr record a negative shift.

The VCD section

In this section, the data obtained in limestone samples for U, V, Mo and As contents show a behaviour similar to that of the Fiume-Bosso section, with the difference that peaks are less pronounced in comparison to the

Table 1 Redox-sensitive trace elements analyses in limestones for the three studied sections (expressed in ppm). The dark grey band indicates the Faraoni Level or its equivalent. No value is indicated when below the limit of detection

Fiume-Bosso		Sample	U	V	Mo	As	Co	Cd	Cu	Zn	Ni	Pb	Cr	
late Hauterivian	no ammonite zonation	FB 59	0.15	0.81		0.39	2.61	0.06	2.61	8.96	13.30	0.90	1.81	
		FB 65	0.05	0.58		0.25	1.51	0.14	3.69	7.21	12.31	1.27	2.13	
		FB 71	0.09	0.64		0.23	3.02	0.12	5.20	7.43	12.24	1.48	2.38	
		FB 79		0.63		0.16	0.99	0.09	2.26	6.01	12.10	1.02	2.09	
		FB 89	0.07	0.71		0.08	1.43	0.14	5.06	9.05	13.76	0.89	2.18	
		FB 94	0.11	0.79		0.09	1.87	0.26	3.11	10.22	13.58	1.23	2.36	
		FB 99	0.08	0.80		0.17	2.06	0.35	4.22	19.28	14.69	5.07	3.19	
		FB 104	0.18	1.25	0.03	0.31	2.26	0.10	6.74	20.41	16.34	2.00	1.85	
		FB 109	0.64	1.10	0.01	0.16	1.70	0.07	3.91	4.64	15.09	6.63	1.62	
		FB 115	0.12	0.91		0.33	1.92	0.14	3.48	8.91	15.30		1.83	
	FB 123	0.15	0.93	0.10	0.12	2.28	0.21	5.24	31.52	15.04		1.70		
	angulicos. / catulloi	FB 128a	0.19	0.91	0.03	0.22	1.64	0.28	2.92	17.02	13.45	2.87	3.10	
		FB 134a	0.24	1.08	0.16	0.51	2.07	0.08	5.70	17.88	15.64	1.53	1.77	
		FB 136	0.23	0.73	0.03	0.46	1.48	0.10	4.78	24.86	13.74	0.91	1.95	
		FB 139	0.39	1.08		0.21	1.23	0.16	2.56	5.23	11.14	1.80	2.01	
		FB 371	0.55	2.04	0.09	0.18	1.08	0.09	2.64	8.95	10.37		1.99	
		FB 373a	0.49	4.60	0.42	0.50	1.36	0.37	1.56	22.37	12.40		1.62	
		FB 373d	0.76	5.21	0.55	0.29	0.96	0.27	1.68	9.14	11.67		1.97	
		FB 375	2.53	11.61	0.31	0.42	0.94	0.19	1.84	16.43	10.12		2.44	
		FB 378	3.10	10.36	0.32	0.23	0.92	0.29	3.27	7.42	11.36		2.16	
FB 380a		1.09	7.74	0.29	0.31	0.89	0.04		13.40	13.22		2.18		
FB 380b	1.80	7.35	0.40	0.27	0.79	0.18		28.41	10.03		1.97			
FB 382a	2.08	6.96	0.15	0.32	1.01	0.09	1.36	7.07	11.37		2.29			
FB 382c	0.81	2.78	0.20	3.73	2.25	0.14		26.70	11.24		1.70			
FB 385	0.64	3.01	0.14	3.37	1.49	0.09	2.16	8.69	11.87		1.84			
FB 387b	0.37	2.08	0.14	0.39	1.79	0.17	3.54	9.88	11.96		3.19			
FB 391	0.24	0.84	0.03	0.29	2.35	0.26	3.01	8.32	11.93		1.79			
FB 396b	0.43	0.82	0.08	0.44	2.02	0.08	1.92	7.97	11.82		1.49			
FB 402b	0.12	0.63	0.12	0.20	1.37	0.14	2.12	20.70	12.44					
FB 410	0.22	1.13	0.16	0.43	2.33	0.13	2.65	10.22	15.59	0.98				
early Barremian	no ammonite zonation	FB 417	0.13	0.89		0.24	1.05	0.11	1.48	5.95	12.97			
		FB 424	0.09	0.65	0.05	0.19	1.37	0.16	2.65	5.80	12.72			
		FB 430a	0.30	0.89		0.26	1.51	0.25	2.32	9.71	13.31	0.83		
		FB 433	0.11	1.00	0.01	0.22	0.97	0.12	1.88	6.88	11.71			
		FB 440a	0.13	1.73	0.02	0.28	1.93	0.18	3.58	12.15	15.39	1.03	1.60	
		FB 448b	0.18	0.86	0.01	0.12	1.11	0.11	2.01	6.24	12.77			
		FB 451a	0.25	0.89	0.04	0.35	1.35	0.10	2.26	5.09	12.06			
		FB 454b	0.13	0.85	0.03	0.12	1.13	0.13	2.14	7.02	12.78			
		FB 459b	0.11	0.68		0.13	1.21	0.13	2.46	5.20	11.97			
		FB 464a	0.14	1.09	0.03	0.37	1.04	0.10	1.93	9.53	7.96			
	FB 469	0.16	0.93	0.04	0.32	1.25	0.10	2.72	6.54	12.49				
	FB 473b	0.29	0.85		0.21	1.14	0.06	1.85	5.73	13.86				
	FB 480a	0.13	0.85	0.10	0.24	1.27	0.10	2.88	17.81	12.31		1.96		
	FB 487	0.10	0.85	0.04	0.29	1.86	0.23	3.18	3.34	17.36				
	FB 494b	0.23	1.08	0.04	0.24	1.22	0.09	1.71	41.15	16.10				
	FB 499a	0.45	1.82	0.17	1.76	1.45	0.11	2.84	66.11	15.79	2.32	1.42		
	late Hauterivian	sayni	VCD 1A	0.64	8.74	0.27	1.38	2.28		7.55	26.45	11.68	2.22	7.73
			VCD 3A	0.52	6.28	0.02	0.45	2.08		7.44	23.04	16.91		6.58
			VCD 4A	0.59	4.92	0.25	0.56	1.79	0.03	8.04	14.88	18.34		4.97
			VCD 6A	0.43	5.02		0.25	1.99		7.26	19.28	17.01		5.01
VCD 9A			0.42	4.36	0.10	0.13	1.37		6.41	12.55	14.31		5.11	
VCD 13A			0.72	6.27	0.33	0.46	1.91	0.01	7.98	15.41	17.41		6.14	
VCD 15A			0.54	6.17	0.08	0.64	2.06		7.40	12.24	16.34		4.59	
VCD 18A			0.77	6.40	0.19	0.46	1.29		7.08	15.07	14.44		4.85	
VCD 21A			0.76	5.82	0.34	0.66	1.76	0.03	9.08	17.76	16.73		6.02	
VCD 24A			0.48	5.79	0.41	0.70	1.64		6.70	13.81	16.87		5.56	
VCD 27A		0.65	8.14	0.21	1.24	1.92	0.04	6.92	19.41	13.19	4.67	5.97		
VCD 29A		0.43	6.98	0.15	0.91	2.12	0.08	9.10	31.52	14.93	2.84	5.92		
ligatus		SDII 1	0.48	6.60	0.29	1.21	2.16	0.07	7.63	20.81	14.79	3.49	5.12	
		SDII 3	0.59	4.09	0.24	0.88	1.77	0.03	6.29	23.44	12.49	2.53	2.88	
		SDII 5	0.52	5.17	0.07	0.82	1.51	0.05	5.43	21.42	20.02	4.22	4.50	
		SDII 6	0.52	4.95	0.15	0.37	1.28	0.04	5.76	17.89	11.12	2.34	4.07	
		SDII 7c	0.45	5.57	0.18	0.36	1.44	0.08	4.70	21.01	11.31	2.27	3.62	
		SDII 10	0.53	6.33	0.19	0.85	1.93	0.03	8.60	23.37	15.30	5.20	4.51	
		SDII 11	0.48	7.77	0.47	0.66	2.25	0.02	6.69	22.48	14.73		5.72	
		SDII 13a	0.40	6.06	0.31	0.67	1.79	0.04	6.33	20.49	12.03	5.22	3.97	
		SDII 14	0.51	4.33	0.15	0.55	1.76		6.25	15.62	12.48		2.43	
		SDII 15	0.53	5.89		0.79	2.12		6.01	15.90	13.55		4.14	
balearis		SDII 16b	0.59	6.38		0.56	1.77		45.52	29.02	12.91		3.53	
		SDII 18	0.54	5.26	0.26	0.67	1.93		5.45	17.30	12.72		4.04	
		SDII 19	0.38	5.66	0.04	0.76	2.43	0.28	6.72	93.62	14.17	6.48	3.42	
		SDII 21	0.68	4.92	0.14	0.47	1.64		4.05	17.02	12.65		4.36	
		SDII 23	0.71	4.76		0.47	1.96		5.12	16.02	13.27		3.78	
		SDII 25	0.57	5.54	0.02	0.62	1.37		9.60	14.59	11.24	1.87	3.76	
		SDII 27	0.40	4.72		0.51	1.68		7.00	19.06	9.74	1.96	4.03	
		SDII 28	0.53	5.32	0.06	1.53	3.94		10.82	31.10	19.45	4.59	4.16	
		SDII 30	0.48	4.64		0.32	1.89		6.18	28.28	10.29	1.81	3.75	
		SDII 33	0.74	7.73	0.12	0.95	2.79	0.07	11.08	37.47	16.66	3.37	8.04	
angulicostata		SDII 34	0.53	5.25		0.88	3.06		5.57	13.85	13.84	2.69	4.61	
		SDII 36	0.53	4.46	0.01	1.28	2.60		5.38	27.48	13.70	3.48	2.30	
		SDII 40	0.61	6.17		0.70	1.47		6.60	22.08	10.30	1.49	5.37	
		VCDII 51A	0.42	4.33	0.10	0.61	2.40	0.09	6.76	22.40	10.58	2.24	4.39	
		SDII 42	0.47	5.82		0.60	1.69		5.42	16.23	10.90	2.10	4.11	
		VCDII 50A	0.64	4.96		0.47	1.35	0.03	6.25	13.85	9.02	1.34	3.57	
		SDII 43	0.64	7.88	0.21	0.53	1.26	0.11	9.11	22.73	11.26	2.24	6.28	
		VCDII 49A	0.60	6.28		1.01	3.11	0.14	7.24	38.48	13.49	1.64	4.93	
		VCDII 48A	0.60	8.57		1.00	2.66	0.17	7.31	47.85	12.39	2.68	6.31	
		VCDII 46A	0.75	5.33		0.45	1.71	0.03	7.64	24.44	10.43	1.25	3.61	
VCDII 44A		0.56	5.88		0.58	1.00		4.70	12.86	8.48		3.79		
VCDII 42A		0.53	5.15		0.54	1.86	0.05	5.89	13.90	8.57	0.89	4.19		
VCDII 40A		0.73	8.04	0.26	1.14	2.32	0.07	7.46	16.68	11.50	1.47	6.77		
SDII 51		0.49	6.52		0.44	0.99		7.19	19.08	10.36	1.47	5.11		
VCDII 37A		0.60	5.06		0.48	1.39		3.93	13.15	8.45	0.86	3.95		
VCDII 35A		0.39	4.82	0.01	0.44	1.32	0.02	4.59	13.49	7.57	1.16	3.06		
VCDII 33A		0.40	4.83		0.57	1.77		5.47	14.92	8.82	1.51	3.25		
VCDII 30A		0.39	4.39		0.48	1.11		4.59	10.09	13.62	0.96	3.33		
VCDII 27A	0.37	5.02		0.68	1.85		3.90	22.87	9.85	0.89	2.66			

Table 1 continued

VCD		Sample	U	V	Mo	As	Co	Cd	Cu	Zn	Ni	Pb	Cr	
late Hauterivian	angulicostata	VCDII 19A	0.59	7.16		1.15	2.34		7.02	21.65	14.25	2.73	5.43	
		VCDII 16A	0.94	11.69	1.22	2.54	1.51		5.44	13.43	11.39	1.13	3.60	
		VCDII12A	0.68	6.79	0.75	1.13	1.66		0.05	5.27	13.92	6.94	1.19	3.28
		VCDII10A	0.53	3.96		0.73	1.45	0.02		5.45	14.87	7.66	1.59	2.40
		VCDII8A	0.46	4.48		0.51	1.57	0.03		6.46	17.02	7.54	1.28	3.81
		SD 42	0.44	6.50		0.75	1.88	0.05		5.53	20.46	12.37	4.58	5.72
		VCDII6A	0.41	4.14		0.74	1.63			4.41	15.89	8.05	1.44	2.15
		SD 38	0.48	6.45	0.31	0.89	1.62		0.04	5.81	17.32	11.00	1.47	3.83
		SD 36	0.43	5.24		0.38	1.22	0.03		6.02	20.50	11.61		2.94
		SD 32	0.41	5.88		0.51	1.12	0.05		4.66	16.93	12.01	1.30	4.07
		SD 29	0.37	5.39	0.07	0.88	1.91			5.25	17.39	10.62	2.20	4.09
		SD 27	0.49	1.88			1.24		0.01	5.65	15.35	12.01	2.43	
		SD 24	0.50	1.79				2.02		5.72	18.61	11.17	1.30	
		SD 21	0.36	2.17						5.69	7.73	11.63	1.61	
SD 19	0.44	2.69				1.25		5.97	11.26	11.74	1.92			
SD 16	0.37	2.54				1.42		5.60	16.38	18.00	2.43			
SD 12	0.49	2.67				1.22		7.72	12.33	11.73	1.44			
SD 8	0.37	1.96				1.99		6.60	18.12	10.74	1.56			
SD 5	0.42	4.14			0.05	2.36		5.82	14.32	12.02	1.84			
SD 1	0.46	5.18			0.30	2.61		6.08	25.27	14.83	3.05			
Barremian	hugii													
Angles		Sample	U	V	Mo	As	Co	Cd	Cu	Zn	Ni	Pb	Cr	
late Hauterivian	balearis	AN 2	0.57	7.78	0.60	0.82	1.04	0.02	5.54	16.12	13.14	1.52	6.69	
		AN 4	0.50	4.88	0.26	0.65	1.36	0.02	4.47	17.42	12.63	2.00	3.87	
		AN 6	0.60	5.04	0.20	0.77	1.49	0.03	4.68	21.93	13.41	2.35	3.68	
		AN 8	0.71	5.03	0.33	0.44	1.01	0.06	5.91	35.59	12.25	1.25	4.79	
		AN 10	0.88	4.87	0.14	0.77	1.79	0.02	5.79	15.46	13.57	1.54	4.44	
		AN 13	0.64	3.64	0.11	0.29	0.97	0.01	3.72	13.18	11.18		3.29	
		AN 15.1	0.57	2.62	0.19	0.39	1.23	0.03	3.58	11.36	11.29	1.05	2.75	
		AN 16	0.52	2.75	0.14	0.36	1.28		3.64	7.65	11.24	1.28	2.33	
		AN 18	0.55	3.89	0.40	0.45	1.32	0.05	4.88	30.37	11.16		3.70	
		AN 20	0.49	3.43	0.45	0.25	0.80	0.04	3.11	17.93	8.95	0.91	3.41	
		AN 23	0.91	4.13	0.16	0.41	1.36	0.07	6.66	17.79	12.45	1.48	3.61	
		AN 26	0.76	2.92	0.18	0.30	0.83		4.32	14.90	9.81	0.98	3.79	
		AN 27.2	0.62	4.64	0.25	0.66	1.15	0.11	6.85	26.72	11.30	1.91	4.43	
		AN 29	0.55	2.70	0.44	0.73	1.11	0.03	3.35	13.45	10.79	1.55	2.63	
		AN 31	0.67	3.04	0.17	0.61	1.17	0.02	4.61	20.26	11.28	1.24	3.18	
		AN 33	0.47	2.71	0.11	0.40	1.06		3.42	6.85	10.55	1.16	3.11	
		AN 35	1.02	4.02	0.13	0.50	1.05	0.09	5.49	16.32	11.66	1.62	2.99	
		AN 38	0.48	2.45	1.02	1.49	5.15	0.39	5.92	75.44	24.32	6.24	2.39	
	AN 40	0.63	4.15	0.57	0.61	0.98	0.20	4.40	23.53	10.43	1.99	3.30		
	AN 43	0.73	3.97	0.24	0.57	1.25	0.09	5.67	18.99	11.65	1.93	3.81		
	AN 46	0.49	2.85	0.06	0.35	1.08	0.04	3.60	15.65	11.06	0.96	2.90		
	AN 49.1	0.73	4.18	0.10	0.43	1.13	0.11	3.93	48.61	10.54	1.13	3.00		
	AN 50	0.49	3.45	0.36	0.35	0.77	0.06	3.22	14.54	9.72	1.21	2.65		
	AN 51	1.05	4.62	0.59	0.74	1.16	0.08	5.34	26.19	12.53	1.58	3.48		
	AN 52	1.43	8.93	4.87	4.80	2.36	2.12	6.41	420.80	26.63	5.73	2.14		
	AN 53.2	1.28	7.43	4.61	3.45	1.42	0.05	3.83	13.13	23.69	4.81	1.99		
	AN 55	2.07	4.75	0.14	0.20	0.71	0.03	3.16	25.79	8.58	1.02	2.35		
	AN 62.1	0.78	4.79	0.44	0.64	1.03		3.52	17.13	10.21	1.08	2.45		
	AN 65.1	0.98	4.38	0.16	0.54	1.56		3.93	12.59	11.89	2.06	1.87		
	AN 68	0.79	4.61	0.20	0.46	0.98	0.06	4.69	24.61	10.37	1.03	3.13		
	AN 71	0.89	4.22	0.06	0.47	0.99		3.70	15.91	10.74	1.22	2.40		
	AN 72	1.17	4.61		0.31	1.05	0.12	4.14	23.89	11.19	1.91	2.22		
	AN 75	0.53	3.57	0.05	0.38	1.03	0.02	3.44	16.25	10.10	1.15	2.76		
	AN 78	0.54	5.00	0.17	0.62	2.17	0.04	4.18	29.54	11.93	1.57	3.40		
	AN 81	0.73	5.30	0.04	0.50	1.30	0.08	4.84	33.86	10.64	1.51	2.83		
AN 84	0.52	4.88	0.18	0.42	1.28	0.04	3.50	26.82	11.05	1.35	2.44			
AN 87	0.59	7.34	0.61	0.88	1.57	0.04	5.15	17.40	12.42	2.05	5.31			
AN 89	0.56	5.17	0.21	0.61	1.25	0.03	3.89	14.02	9.25	1.26	3.48			
AN 90	0.56	4.98	0.29	0.86	1.79		4.52	18.55	13.17	2.06	3.59			
AN 92	0.40	4.08	0.20	0.58	1.34	0.04	4.55	12.24	11.60	2.37	3.18			
AN 93	0.37	3.59	0.03	0.46	1.25	0.02	3.62	16.10	11.02	1.36	3.52			
AN 95	0.51	3.95	0.21	0.87	1.47	0.06	5.80	17.18	10.65	1.74	3.31			
AN 97	0.49	3.13	0.07	0.42	1.11	0.06	3.85	15.57	11.56	1.26	2.64			
AN 98	0.53	3.49	0.21	0.75	1.50	0.11	11.51	29.09	10.91	1.40	2.71			
AN 101	0.71	4.92	0.23	0.92	1.33	0.04	6.11	23.13	12.82	2.21	3.65			
AN 103	0.57	6.88	0.65	1.02	2.01	0.06	6.03	30.35	12.74	2.42	4.53			
AN 104	0.52	5.15	0.37	1.10	1.30	0.06	6.05	24.55	11.33	2.22	4.06			
AN 105	0.52	4.47	0.63	1.22	1.49	0.05	6.94	33.44	11.80	2.28	3.79			
AN 108	0.89	5.81	0.71	1.08	1.49	0.08	6.21	37.61	11.22	2.20	3.37			
AN 109.4	0.49	4.14	0.69	1.14	1.72	0.02	5.75	29.79	11.49	2.08	3.04			
AN 110.1b	0.68	5.48	1.11	1.36	1.73	0.04	5.34	52.45	13.14	2.66	3.08			
AN 110.3c	0.62	3.99	0.33	0.67	1.38	0.08	4.04	47.02	10.42	0.91	2.64			
AN 110.4	1.04	5.03	0.47	1.07	1.09	0.07	5.29	21.30	11.21	1.77	2.88			
AN 111.2	0.80	5.36	0.12	0.87	0.88	0.05	4.58	20.12	8.83	1.50	2.58			
AN 111.5	1.22	7.64	0.72	1.12	1.20	0.09	5.10	65.66	10.84	1.86	3.26			
AN 112.2	0.74	6.66	0.17	0.55	1.33	0.07	4.41	23.15	10.43	1.24	2.52			
AN 112.7	0.59	5.55	0.42	0.63	1.04	0.08	4.12	150.20	10.55	1.83	2.41			
AN 114.2	0.68	4.61	0.41	0.88	0.93	0.04	4.68	24.82	10.13	1.43	1.80			
AN 115	0.77	6.43	0.62	1.21	1.05	0.03	5.45	24.46	13.81	9.36	2.57			
AN 117	0.60	4.16	0.30	0.74	1.27	0.05	5.44	24.94	10.40	1.36	1.87			
AN 119	0.57	3.10	0.23	0.34	1.00	1.25	4.83	80.29	10.27		1.55			
AN 121	0.54	3.38	0.12	0.32	1.34	0.06	3.29	18.69	10.94	0.99	1.86			
AN 123	0.57	4.26	0.27	0.53	1.38	0.08	4.57	30.62	11.74	1.10	2.53			
AN 125	0.96	6.99	0.73	1.06	1.95	0.05	7.09	41.84	14.53	1.94	3.81			
AN 127	1.02	4.27	0.32	0.36	0.76	0.03	5.15	18.56	8.52		1.98			
AN 128	0.76	3.24	0.34	0.42	0.84	0.07	4.39	14.86	9.74		1.69			
AN 129	0.75	4.25	1.70	2.58	1.76	0.09	6.48	23.92	12.89	1.88	2.28			
AN 131	0.84	8.34	0.73	0.95	1.25	0.03	7.52	21.95	13.50	1.47	1.90			
AN 133	0.84	3.10	0.39	1.08	1.00	0.10	6.30	21.42	12.97	1.45	2.01			
AN 136.2	0.61	4.88	0.47	1.94	2.27	0.07	8.46	29.82	14.45	3.71	2.89			
AN 138	0.70	3.41	0.57	1.10	0.97	0.21	6.72	44.43	11.66	1.36	1.17			
AN 139d	1.08	3.54	0.16	0.63	0.94	0.17	4.86	26.74	11.87	1.95	3.74			
early Barremian	nicklesi	AN 95	0.51	3.95	0.21	0.87	1.47	0.06	5.80	17.18	10.65	1.74	3.31	
		AN 97	0.49	3.13	0.07	0.42	1.11							

Table 2 Redox-sensitive trace-element analyses in marlstones for the three studied sections (expressed in ppm). The dark grey band indicates the Faraoni Level or its equivalent. No value is indicated when below the limit of detection

Fiume-Bosso	Sample	U	V	Mo	As	Co	Cd	Cu	Zn	Ni	Pb	Cr	
late Hauterivian	angu.	FB 133	2.01	69.51	2.96	15.55	39.00	1.62	93.32	287.33	136.27	39.08	38.66
		FB 138	3.97	27.74	1.00	19.77	41.14	1.20	139.33	235.05	99.30	26.68	33.83
		FB 372	9.47	84.72	1.90	14.29	58.25	3.38	214.15	528.47	223.99	50.01	53.08
	catulloi	FB 374	12.83	345.94	61.83	177.42	84.54	10.48	270.70	1068.19	533.54	49.22	66.55
		FB 377	6.29	68.15	15.97	8.13	9.42	4.32	23.68	220.95	99.15	4.49	8.79
		FB 379	15.40	306.71	19.28	28.51	67.00	4.23	273.71	701.05	428.86	33.57	71.43
		FB 381b	9.92	151.09	28.26	33.44	29.64	5.06	61.67	313.80	268.18	14.52	22.21
		FB 384	12.04	169.53	5.48	40.58	106.59	3.83	211.07	612.52	277.61	51.44	66.48
		FB 397	2.85	20.43	0.30	8.25	67.51	0.90	164.07	248.47	143.25	34.58	33.69

VCD	Sample	U	V	Mo	As	Co	Cd	Cu	Zn	Ni	Pb	Cr	
late Hauterivian	angulicostata	VCDII 10b	1.38	12.59	0.45	1.89	4.14	0.13	14.62	50.26	19.43	2.25	12.93
		VCDII 11b	1.29	8.27	0.29	1.87	3.84	0.21	13.77	49.12	21.39	3.79	8.43
		VCDII 14b	0.82	5.47	0.57	1.95	3.00	0.05	10.19	20.06	19.01	1.90	7.22
		VCDII 17b	3.27	15.21	1.59	6.74	3.46	0.60	15.77	53.27	23.53	3.95	8.34
		VCDII 18b	1.26	11.67	0.12	2.37	4.93	0.05	16.14	51.92	27.56	4.30	13.63
		VCDII 20b	1.45	11.60		3.44	7.42	0.20	18.30	86.81	37.34	9.23	11.95
		VCDII 25b	1.08	8.15		1.85	5.13	0.43	19.39	112.65	28.29	4.37	9.33
		VCDII 26b	1.05	8.56		2.68	4.41	0.13	15.84	56.10	23.71	5.68	9.18
		VCDII 27b	1.31	11.47	0.21	1.94	4.38	0.07	20.00	49.23	28.48	3.75	13.25
		VCDII 28b	1.06	13.37	0.15	2.16	5.01	0.16	16.50	61.06	21.64	4.25	12.78
		VCDII 29b	0.91	7.75	0.06	1.47	3.00	0.13	14.43	36.53	20.36	3.41	7.89
		VCDII 30b	1.03	6.84	0.03	1.44	4.01	0.07	17.12	41.84	22.74	3.64	8.05
		VCDII 31b	1.17	6.99	0.19	2.09	5.13	0.06	15.28	36.51	26.18	5.04	8.26
		VCDII 32b	1.24	9.64	0.03	1.52	4.43	0.17	19.81	47.95	28.04	3.25	13.09
		VCDII 33b	1.50	11.92	0.17	2.30	4.63	0.10	24.18	65.32	34.76	6.90	13.59

Angles	Sample	U	V	Mo	As	Co	Cd	Cu	Zn	Ni	Pb	Cr	
late Hauterivian	angulicostatum	AN37b	1.92	21.45	0.57	2.86	5.76	0.56	42.80	100.62	50.60	12.67	26.37
		AN38b	1.59	14.64	0.44	3.13	5.71	0.38	38.98	81.97	38.78	13.09	16.89
		AN39b	1.62	16.99	1.15	4.12	6.64	1.28	36.73	167.87	42.80	14.17	16.63
		AN40b	2.17	16.34	0.71	3.86	6.89	0.56	44.88	139.18	42.11	22.05	16.67
		AN41b	2.56	22.10	0.73	4.47	8.67	0.84	58.35	205.46	55.25	20.63	23.53
		AN42b	2.28	15.30	0.16	2.99	5.63	0.58	39.87	110.99	42.70	10.33	17.90
		AN43b	1.59	12.39	0.27	1.73	3.39	0.35	24.57	74.26	29.27	7.90	12.99
		AN44b	1.58	12.78	0.28	2.41	3.88	0.34	28.36	72.99	27.64	10.51	14.37
		AN45b	2.32	31.12	0.57	4.80	13.38	0.32	57.41	111.72	55.01	16.73	30.66
		AN46b	2.15	12.78	0.25	3.45	7.80	0.39	48.73	103.01	68.49	9.22	15.47
		AN48b	1.85	18.98	0.37	3.35	6.26	0.49	44.40	114.71	37.86	10.54	21.80
		AN51b	3.54	19.23	2.97	2.06	1.82	0.06	8.39	31.71	33.60	3.61	8.42
		AN52b	3.73	27.35	1.78	4.71	4.19	0.45	19.08	91.12	51.41	7.13	17.08
		AN53.1b	2.98	18.84	3.30	1.52	1.69	0.06	10.16	19.16	38.32	3.13	7.90
Barr.	kilian.	AN72b	1.59	10.45	0.37	3.20	3.72	0.22	25.72	136.80	24.11	7.30	8.84
		AN73b	1.61	10.84	0.31	2.94	6.00	0.50	22.91	138.40	32.25	7.44	10.76
		AN75b	1.48	13.92	0.19	3.39	5.66	0.29	26.22	128.48	33.94	8.73	14.48

V present a double peak in contents around the level of the “couches à poissons,” with the higher peak representing the “couches à poissons” itself. The average values of the background levels correspond to approximately 0.5, 5.4, 0.09 and 0.8 ppm and the maximum values are approximately 0.9, 11.7, 1.2 and 2.5 ppm, for U, V, Mo and As contents, respectively. The data obtained in the marls show equally a clear enrichment of U, V, Mo and As around the Faraoni Level (Table 2).

Co contents in limestone show a small overall decrease from the bottom to the top of the section, with a mean value of about 1.7 ppm. A peak is

distinguished 5 m below the equivalent of the Faraoni Level, reaching 3.6 ppm. In the marls, a peak reaching 7.4 ppm is distinguished 3 m below the Faraoni Level.

In this section, the Cu, Zn, Ni and Pb show a small positive shift in the limestone 5 m below the Faraoni Level equivalent. They correspond to the peaks of U, V and Co, Cu and Zn show equally a peak approximately 40 m below the Faraoni Level, at the beginning of the *B. balearis* zone. On the other hand, in the marls, only Cd is enriched within the Faraoni Level equivalent, whereas enrichment of Cd, Zn, Ni and Pb characterizes again the beds situated 5 m below this level.

The Angles section

In the limestone samples, the trends in U, V, Mo and As contents are comparable to the other two sections, with peak values at the level of the Faraoni Level equivalent (Fig. 5; Table 1). Whereas the background levels are low and rather constant for U, Mo and As, V contents show a more variable evolution, with a series of positive excursions above the Faraoni Level equivalent. The average values for the background levels are at about 0.7, 4.5, 0.2 and 0.6 ppm and the maximum values correspond to approximately 2.1, 8.9, 4.9 and 6.3 ppm, for U, V, Mo and As contents, respectively. The data obtained in the marl samples show equally a clear enrichment of U, V, Mo and As around the Faraoni Level equivalent (Table 2).

The overall trend in Co contents in limestone is more or less constant, with average values oscillating around 1.3 ppm. A peak with a value of 5.1 ppm is distinguished approximately 6 m below the equivalent of the Faraoni Level (bed AN 38). In the same bed, a small peak is present in the Mo and the As curves. Within the equivalent of the Faraoni Level, a small peak of Co reaching 2.4 ppm is distinguished. The data obtained in the marls show equally a clear enrichment of U, V, Mo and As around the Faraoni Level equivalent (Table 2). In the marl samples, a remarkable peak in Co contents characterizing the Faraoni Level equivalent has not been observed. A maximum of 13.4 ppm is seen within bed AN 45b.

In this section, a clear peak of Cd, Zn, Ni and Pb occurs within the limestone beds of the Faraoni Level equivalent. Moreover, a peak of Ni and Pb marks bed AN 38. On the other hand, no significant peak of these elements has been found in the Faraoni Level equivalent marls. We notice, however, that all these elements are more concentrated in the marls below the Faraoni equivalent than above.

Total organic carbon

In the limestone samples, measured TOC values are not higher than 0.1%. We will therefore restrict our discussion to the results obtained in the marl samples (see Fig. 6). In the Fiume-Bosso section, a maximum TOC value of 14.8% has been measured in the thin marl level interval just below the Faraoni Level. Within the Faraoni Level, the average value of TOC is close to 3.25%. These values are not significantly different from those obtained in marl samples above and below the Faraoni Level.

In the VCD section, no major variations have been measured (Fig. 6). The average value is close to

Table 3 Background level values, maximum value of the peak and enrichment factors of U, V, Mo and As for the three studied sections (see text for calculation details)

Ppm	Background level average value	Maximum value of the peak	Enrichment factors
Fiume-Bosso			
U	0.17	3.1	18.24
V	0.9	11.6	12.89
Mo	0.03	0.55	18.33
As	0.3	3	10.00
Veveyse			
U	0.53	0.94	1.77
V	5.411	11.7	2.16
Mo	0.09	1.22	13.56
As	0.75	2.54	3.39
Angles			
U	0.69	2.1	3.04
V	4.55	8.9	1.96
Mo	0.24	4.87	20.29
As	0.65	4.8	7.38

0.5% and the Faraoni Level equivalent TOC value is equal to 0.31%. In the Angles section, no major variation has been observed as well. The TOC values are however higher and reach 2.63%. TOC values in the Faraoni Level equivalent are close to 1.25%.

Discussion

Comparison between the three sections

Trends in U, V, Mo and As contents are all characterized by pronounced positive peaks marking the Faraoni Level and its equivalent, and are very distinct relative to the measured intervals below and above this level, representing the upper Hauterivian and the lower Barremian. This positive peak is present in both limestone and marl samples, despite differing background level values. For the trends in Co, Cd, Cu, Zn, Ni, Pb and Cr contents, no significant overall variations are observed and no specific enrichments are recorded for the Faraoni Level and its equivalent apart from local variations.

Behaviour of U, V, Mo, Co and As

Enrichments in U, V and Mo are well known for numerous black shales of the Phanerozoic (e.g. Brumsack and Gieskes 1983; Brumsack 1989, 2006; Coveney et al. 1991; Sun and Püttmann 1997; Nijenhuis et al. 1999; Warning and Brumsack 2000; Yarinçik

Fig. 3 Redox-sensitive trace elements distributions for the Fiume-Bosso section. The grey band indicates the position of the Faraoni Level and the dashed line the limit of quantification for each element

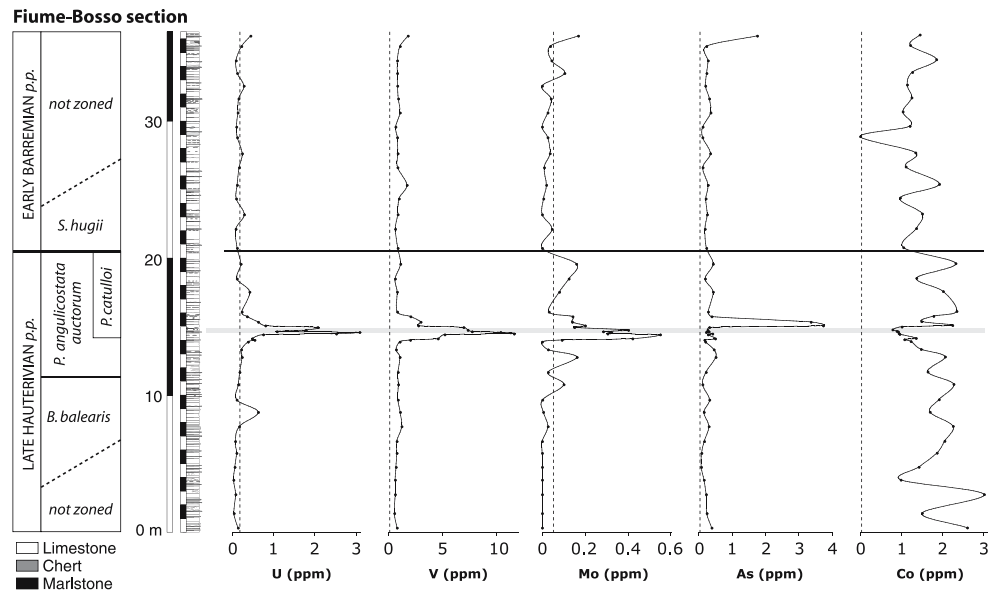
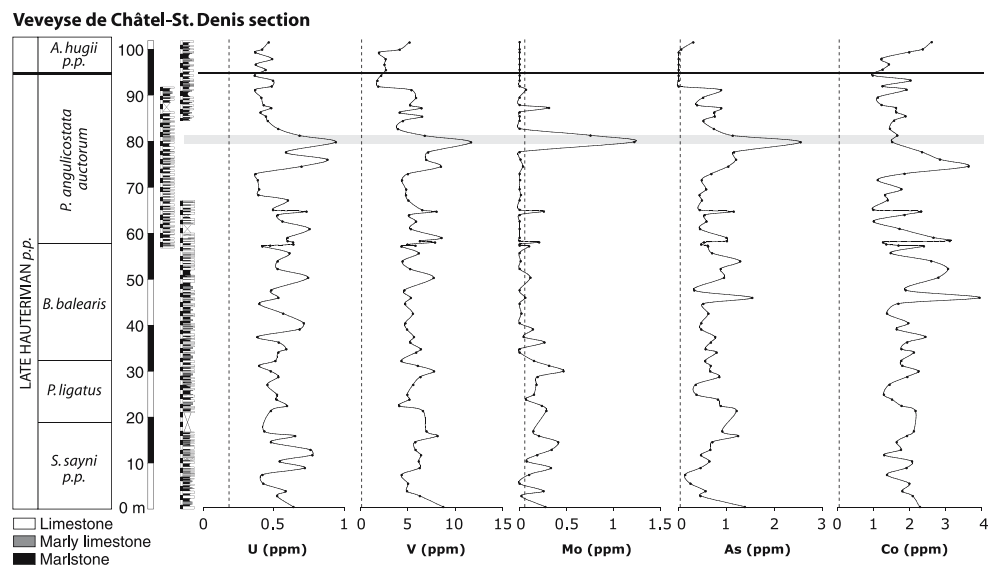


Fig. 4 Redox-sensitive trace elements distributions for the Veveyse de Châtel-St Denis section. The grey band indicates the position of the “couche à poissons” (Faraoni Level equivalent) and the dashed line the limit of quantification for each element



et al. 2000; Mangini et al. 2001; Tribouillard et al. 2004; Wilde et al. 2004). U and V enrichments are usually associated to anoxic to euxinic bottom-water conditions during the time of deposition (e.g. Hastings et al. 1996; Morford and Emerson 1999; Mangini et al. 2001). Under mildly reducing and anoxic conditions, U and V may be removed from the ocean by the formation of organo-metal ligands in humic acids; by crystalline precipitation [UO_2 and V_2O_3 or $V(OH)_3$], and V may also be removed by surface-adsorption processes (Algeo and Maynard 2004). For Mo, Algeo and Maynard (2004) propose adsorption onto humic substances as a main possible mechanism to transfer this element from

the water column into the sediment, and under euxinic conditions, by a rapid uptake by authigenic/syngenetic sulfides.

Overall, the solubility of these three elements decreases in the reduced state, and their precipitation and enrichment in sediments is favored during periods of anoxia (e.g. Yarincik et al. 2000; Mangini et al. 2001; Algeo and Maynard 2004).

Co shows a more complex behaviour in the three sections studied here. This element should behave similar to U, V and Mo, and indeed, in oxygen-depleted waters, Co forms an insoluble sulfide (CoS) that is trapped in sediments (Algeo and Maynard

Fig. 5 Redox-sensitive trace elements distributions for the Angles section. The *grey band* indicates the position of the Faraoni Level equivalent and the *dashed line* the limit of quantification for each element

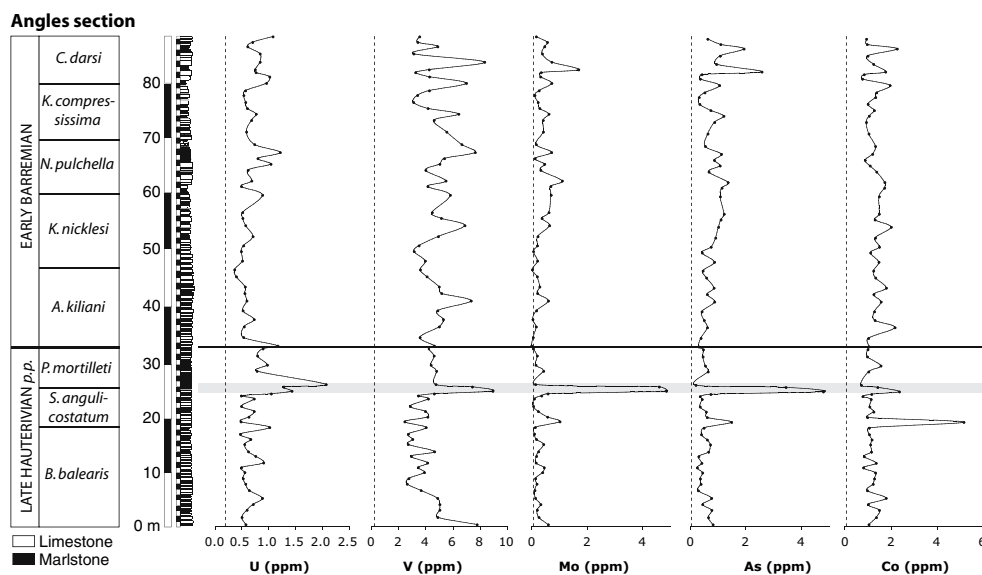
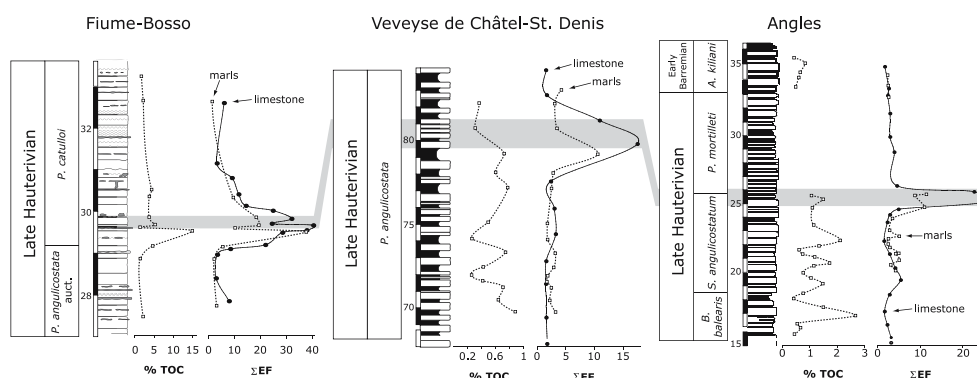


Fig. 6 Correlation of the Faraoni Level for the three different sections. For each section, the total organic carbon analyses in marls as well as the trace element index (ΣEF) in both limestone and marls are figured. See text for ΣEF calculation details. The lithologic legend for each section is the same as in previous figures



2004). These sulfides are taken up in solid solution by authigenic Fe sulfides (Huerta-Diaz and Morse 1992). But, because of the slow kinetics of this uptake, significant Co enrichment in the sedimentary environment is limited to euxinic conditions (Sun and Püttmann 1997; Algeo and Maynard 2004).

Sedimentary enrichments in As are less well documented from the geological record. As is usually mobile under slightly reducing conditions through the reductive dissolution of the As bearing phases [Fe (hydr)oxides]. Furthermore, Bostick and Fendorf (2003) detailed the mechanisms of As uptake in more anoxic environments and demonstrated the surface precipitation of As on iron sulphide minerals. This hypothesis is supported by Huerta-Diaz and Morse (1992) who noted a close link between As and pyrite in sediments underneath anoxic water. Hence, the presence of abundant marcasite nodules within the Faraoni Level in central Italy (Cecca et al. 1994) is compatible with the enrichment in As.

An increased input of trace metals in the ocean in association with large igneous province (LIP) activity has been forwarded as an alternative explanation for trace-metal enrichment in mid-Cretaceous organic-rich sediments, which were formed during oceanic anoxic events (e.g. Sinton and Duncan 1997; Larson and Erba 1999; Leckie et al. 2002; Erba 2004; Snow et al. 2005). For the late Hauterivian, however, there is no evidence for LIP volcanism so far (e.g. Larson 1991; Courtillot and Renne 2003), although a certain degree of volcanic activity has been documented at Rio Grande, documenting the end of the Paraná continental flood basalts (Stewart et al. 1996). The volume of lava emitted during this episode is one order of magnitude lower compared to other LIPs. Moreover, Brumsack (2006) pointed out that trace-metal enrichment (with the exception of extreme enrichment of Ag, Cu, Ni, V and Zn) in sediments associated with the Cenomanian–Turonian boundary oceanic anoxic event may be explained by seawater as the mere source for these elements (see also Arthur et al. 1988; Piper 1994; Böning

et al. 2004). The trace-metal enrichments calculated for the Faraoni event (Table 3) are on the same order of magnitude as those calculated for present-day coastal upwelling zones and anoxic basins (e.g. Brumsack 2006). We therefore exclude volcanic activity as a major source of trace-metal enrichment during the Faraoni event.

Redox conditions during the Faraoni event

The behaviour of the redox-sensitive elements, measured in and around the Faraoni Level and its equivalents, allows us to precise the redox conditions that were present during the Faraoni event. The enrichment of U, V, Mo and As are interpreted as a signature of anoxic and euxinic conditions (Algeo and Maynard 2004). The conservative behaviour of Co during the Faraoni event with regard to its background level is characteristic of dysoxic to anoxic facies according to these authors. The Faraoni Level is therefore clearly the expression of an anoxic event (with eventually locally weakly euxinic conditions such as is indicated at Angles with regard to a small increase of the Co content). The fact that the other redox-sensitive trace elements studied here are not systematically enriched within the Faraoni Level underlines the limit of our comprehension of this system. Partial diagenetic remobilization or differences in sediment uptake processes may be evoked to explain the non-systematic presence of enrichments in the other studied trace elements. Such remobilization processes are thus thought to be responsible of the slight shift of Mo and As peak within the Angles section. Further investigations are however needed to resolve this problem.

Using total sulphur versus TOC plots, Baudin et al. (2002) argued for dysoxic rather than for truly anoxic conditions for the Faraoni Level. This observation together with the conservative behaviour of Co may hint at the possibility that the anoxic conditions were not so strong and closer to the dysoxia/anoxia threshold than to the anoxia/euxinic threshold (O_2 close to 0.2 ml/l of H_2O), with the exception of the section at Angles, where conditions may have been more severe.

Local oxygen-deficient conditions below the Faraoni Level

In the Angles and the VCD sections, isolated maxima in trace-metal contents are recognized, notably below the Faraoni Level equivalent. In the VCD section, U, V, Zn, Ni and Co contents show a peak 5 m below the Faraoni equivalent. In the Angles section, Mo, As, Ni,

Pb and Co contents show a peak 6 m below the Faraoni Level equivalent. These peaks are not correlated in time between the two sections and could have resulted from local changes in bottom-water oxygen conditions, resulting in local anoxia or euxinia.

Enrichment factors

For each separated lithology (i.e. limestones or marls), enrichment factors (EFs) were calculated for U, V, Mo and As according to the following equation (see Table 3; Fig. 6):

$$EFs = [Me]_{\text{peak}}/[Me]_{\text{background}} \quad (1)$$

In this equation, $[Me]_{\text{background}}$ represents the average value of the background level of the different curves and $[Me]_{\text{peak}}$ is the maximum value of the different peaks. The EFs thus obtained are coherent with those obtained by Morford and Emerson (1999) (for U, V and Mo, and based on references therein) for anoxic conditions, representing another clue for anoxic conditions during the Faraoni event.

In order to compare the different sections, an enrichment factor index ($\sum EF$) is calculated by adding the EFs of U, V and Mo (see Fig. 6). A clear peak is thus obtained for the three sections for both lithologies, underlining the higher uptake of U, V and Mo in the sediment during the Faraoni event.

A remarkable feature of the Faraoni Level is the alternation between limestones and laminated shales, which is interpreted, in a first approximation, as the results of variations in oxygen content (see Westphal et al. 2004 for a review). The higher values of redox-sensitive trace elements in the Faraoni shales compared with the Faraoni limestones seem to confirm this hypothesis. However, if we look at the enrichment factors index ($\sum EF$) calculated for both marls and limestones (Fig. 6), such a dichotomy is not found. The variation of these indices is more or less the same within the same section. On the contrary, it seems that the limestones are a little bit more enriched in redox-sensitive trace elements than the marls during the Faraoni event compared to “normal” conditions. The higher redox-sensitive trace element content of the marlstones (compared to the limestones) is thus not induced by difference of synsedimentary redox conditions between marlstones and limestones (as already shown by Westphal et al. 2004), but is rather the result of synsedimentary condensation or early diagenetic processes. This underlines the importance of the lithology effect on analyses and the need to consider

“mono-lithologic” analyses when discussing temporal variations.

In Fig. 6, where the TOC is plotted versus the $\sum\text{EF}$, we remark that TOC is only clearly enriched in the Faraoni Level of the Fiume-Bosso section, whereas $\sum\text{EF}$ shows a positive shift in all three sections. This allows us to postulate that the enrichment of redox-sensitive trace elements is a more conservative tracer of anoxic conditions than TOC contents. This latter may be more easily removed during diagenetic processes.

Chemostratigraphic tools

The distinct enrichment of the redox-sensitive elements U, V, Mo and As in the sedimentary environment during the Faraoni event appears to be unique for the examined time interval of the Late Hauterivian–Early Barremian and is therefore useful as a tracer of the Faraoni event within the Tethyan realm.

Outside the Tethyan realm, a distinct enrichment in U contents occurs in sediments of the Late Hauterivian of the Izu-Mariana margin, northwestern Pacific, in ODP hole 1149b (ODP leg 185; Plank et al. 2000, upper part of the lithologic unit IV, interval 185-1149B-19R-1; see Figs. F27, F72). This peak is associated with a layer containing elevated quantities of euhedral barite, along with dolomite rhombs and phosphatic fish remains. Plank et al. (2000) argue that this peculiar mineral assemblage indicates an episode of enhanced organic-matter accumulation, which they correlate with the Faraoni event. According to the data presented here, the enrichment in U represents a further strong argument in favor of this assumption. This suggests that the Faraoni event was not restricted to the western Tethyan realm, but may also be recorded in sediments of the northwestern Pacific. Further investigations—especially in sediments of the southern hemisphere—are required to confirm the global character of this event.

In the Fiume-Bosso section, the enrichment in U, V and Mo contents began slightly before and ended slightly after the Faraoni Level s.s. This may indicate that the Faraoni anoxic event may have started slightly earlier and lasted longer than the time associated with the formation of the Faraoni Level itself. This would signify that the organic-rich facies of the Faraoni Level in the Italian sections does not reflect the entire period of oxygen depletion, but rather the time of important organic matter preservation and/or accumulation during the culmination of this event. This may be taken into account when tracing the equivalent of the Faraoni Level by using trace-metal enrichment in sections

were ammonites are absent or extremely rare. The here described trace-metal enrichments trace the Faraoni anoxic event rather than the Faraoni Level itself.

Initiating factors of the Faraoni event

As recognized by Baudin et al. (2002) and Baudin (2005), the Faraoni Level and its equivalent can be qualified as an anoxic event. An important difference between the Faraoni event and other Cretaceous OAE's consists, however, in the fact that the anoxic event is not mirrored by a major shift in the carbon-isotope record (e.g. Erba et al. 1999; Van de Schootbrugge et al. 2000; Company et al. 2005; Godet et al. 2006), but only by a small positive excursion (0.5‰). Godet et al. (2006) explain this attenuated signal by a buffering effect on the carbon-isotope record exerted by the large size of the oceanic dissolved inorganic carbon reservoir. Furthermore, the classical scenario invoking an episode of intensified volcanism and reinforced greenhouse conditions (e.g. Jenkyns 1999, 2003) is difficult to apply to this peculiar event. As previously said, there is no evidence for LIP volcanism during the late Hauterivian so far, even if a certain degree of volcanism has been documented in Rio Grande. This volcanic activity is however one order of magnitude lower than that associated with LIP (e.g. Larson 1991; Courtillot and Renne 2003). We therefore consider the Rio Grande volcanism as not responsible for the Faraoni event, even if slightly stronger greenhouse conditions may have been induced by this volcanic episode.

As discussed by Baudin et al. (2002), the Faraoni Level corresponds to the maximum flooding surface (mfs) of the Ha6 sequence (Haq et al. 1987; Hardenbol et al. 1998). For the VCD section, Busnardo et al. (2003) attributed the “couches à poissons” to the transgressive surface of the Ha6 sequence, but A. Strasser (Fribourg, personal communication) correlates the same horizon also to the mfs of the Ha6 sequence (see also Bodin et al. 2006). The Ha6 mfs corresponds to the maximum of a second-order transgression (Haq et al. 1987; Hardenbol et al. 1998), which led to the connection of the Tethyan and the Boreal realms, as is suggested by belemnite migration patterns observed by Mutterlose and Bornemann (2000) and the presence of boreal nanoplankton in the Angles section around the Faraoni Level (Godet et al. 2006). This second-order transgression may have led to the flooding of large epicontinental areas. Bodin et al. (2006) pointed out the role of this mechanism as the main trigger for the

onset of anoxic conditions during the Faraoni anoxic event, by increased nutrient delivery, and the related increase in primary productivity (see also Company et al. 2005). Baudin et al. (2002) also proposed increased primary productivity as the prime factor leading to the origin of the Faraoni event along the line of the scenario proposed here.

Conclusions

Redox-sensitive trace elemental analysis is a fruitful method to better understand and constrain anoxic events. Indeed, the multi-element analysis discussed here provides additional insight into the oxygenation conditions during the Faraoni anoxic event and serves as a promising correlation tool, especially if biostratigraphic information is lacking.

By comparing the behaviour of the redox-sensitive trace elements with the specific enrichments in U, V and Mo contents, the Faraoni Level appears to be the expression of anoxic conditions at least near the sediment–water interface. This interpretation is supported by the enrichment of As, which is largely controlled by its sorption onto sulfide minerals in anoxic environments (Bostick and Fendorf 2003). The lack of systematic enrichment of Cd, Cu, Zn, Ni, Pb and Cr may be related to diagenetic removal or unknown synsedimentary differences in uptake processes.

In the sections analysed, the distinct enrichments in U, V and Mo are a feature characteristic of the Faraoni Level, and are unique with regards to the time interval of the late Hauterivian and the early Barremian. We therefore propose to use this feature to correlate the Faraoni Level throughout the entire Tethys, and especially in sections where organic-rich intervals are lacking such as in the section of Angles, and to search for expressions of the Faraoni event also in other basins. For example, a distinct enrichment in U in sediments of late Hauterivian age in the northwestern Pacific may well be correlated with the Faraoni event (Plank et al. 2000; see Figs. F27, F72).

According to the general behaviour of the redox-sensitive trace elements in the three analysed sections, the Faraoni Level may be the expression of a climax situation during a longer period of steadily increasing anoxia induced by a second-order sea level rise. The mobilization of nutrients during this sea level rise may have played an important role in the onset of the Faraoni event.

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